

Stochastic Late Accretion on the Earth, Moon and Mars

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Why Study the Moon?



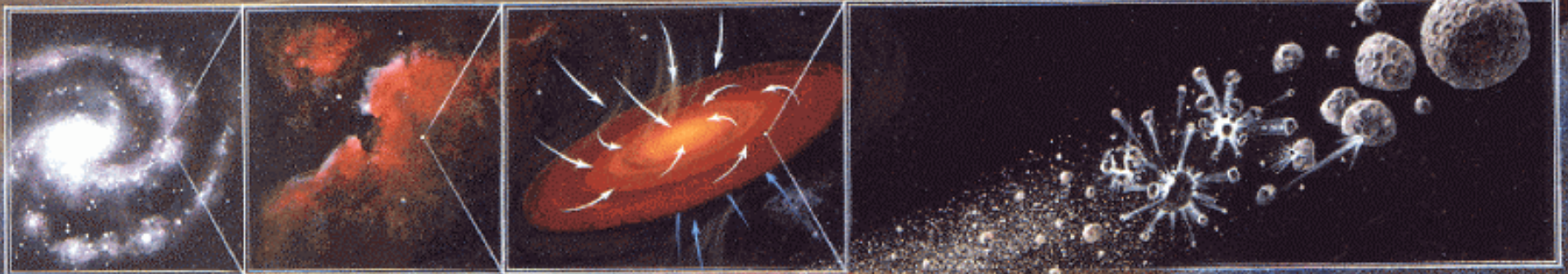
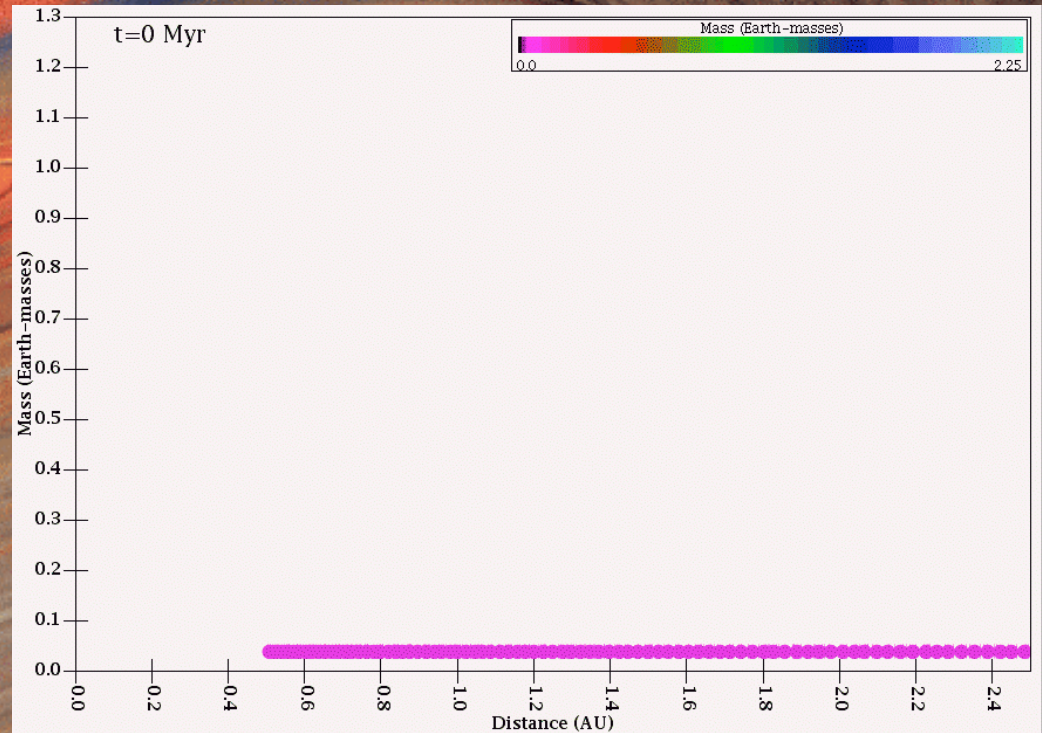
- The Moon itself is fascinating, but it is also a “Rosetta Stone” for telling us about:
 - The unknown nature of the primordial Earth!
 - The critical last stages of planet formation throughout the solar system!

Part 1: **Standard Model of Planet Formation**

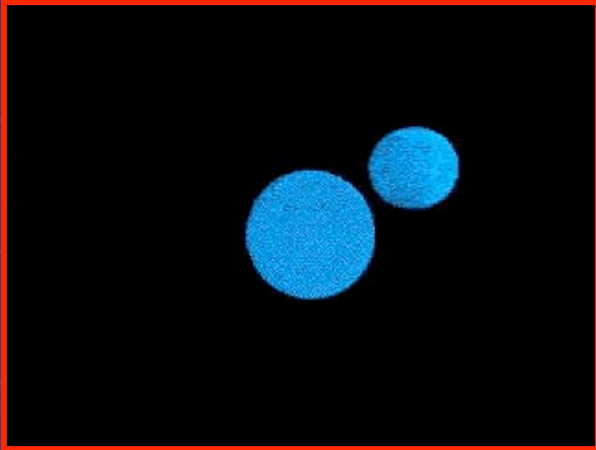


Planetesimal and Planet Formation

- Disk particles come together by gravity.
- Collisions make larger objects by “accretion”.
- Planetary embryos collide and eventually create planets.



Outcomes from the Moon-Forming Impact



- Giant impact leads to last differentiation event on the Earth and Moon near ~ 60 (+90, -10) My after the formation of CAIs at 4.56 Ga.
- Final phase of core formation and global magma oceans occur on both worlds.
- A thick stable lunar crust grows over time. The Earth also grows a crust that can be recycled by plate tectonics.

That's It, Right?



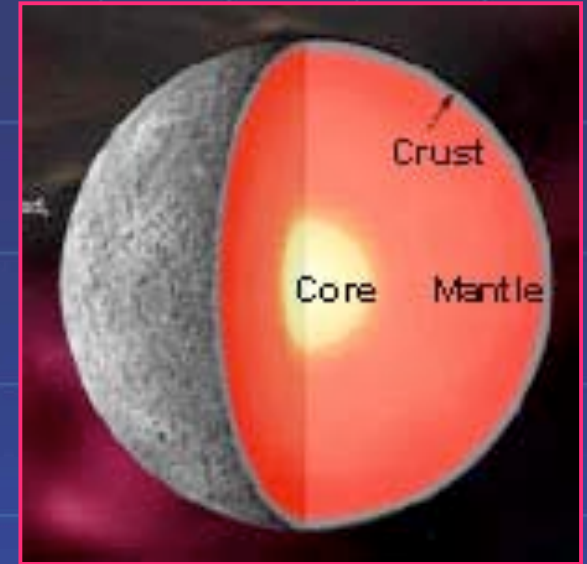
- Are the Earth and Moon effectively done in terms of their internal structure being influenced by impacts?

Part 2:

The Curious Case of the Highly Siderophile Elements

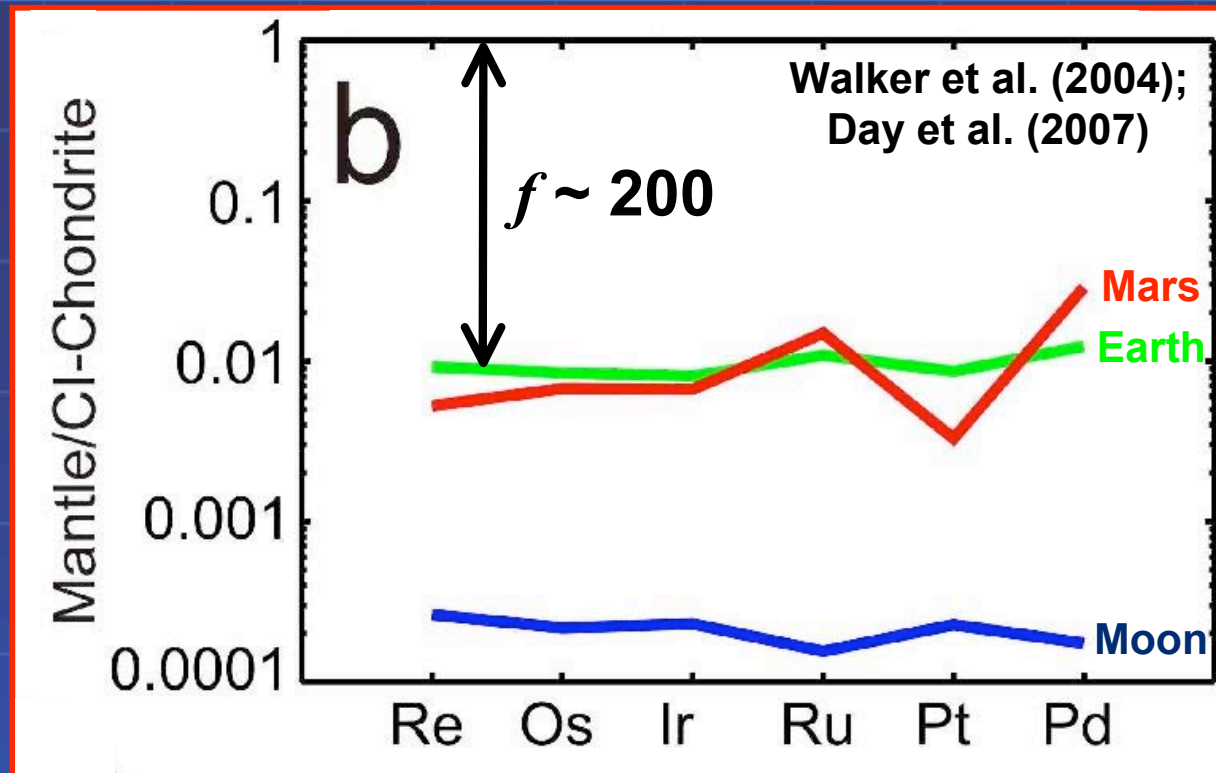


Highly Siderophile Elements (HSEs)



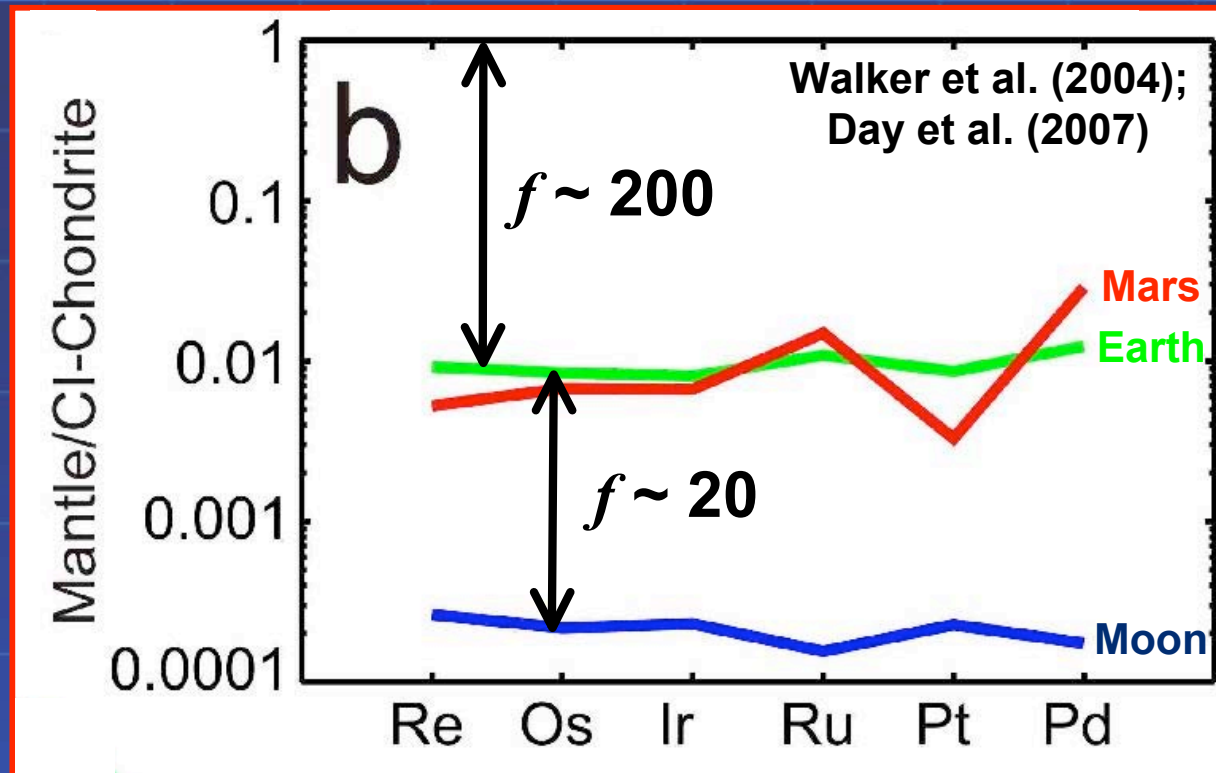
- HSEs (Re, Os, Ir, Ru, Pt, Rh, Au) are metals with high metal-silicate partition coefficients ($> 10^4$).
- During primary accretion, differentiation, and core segregation, HSEs should go to a planet's core, never to be seen again.

Mantle HSEs from Earth, Moon, and Mars



- Earth's ancient mantle only depleted in HSEs by factor of ~200 compared to chondrites. **Why?**
- It also had chondritic relative proportions (i.e., it is pretty “flat” compared to standard chondrite abundances).

Mantle HSEs from Earth, Moon, and Mars



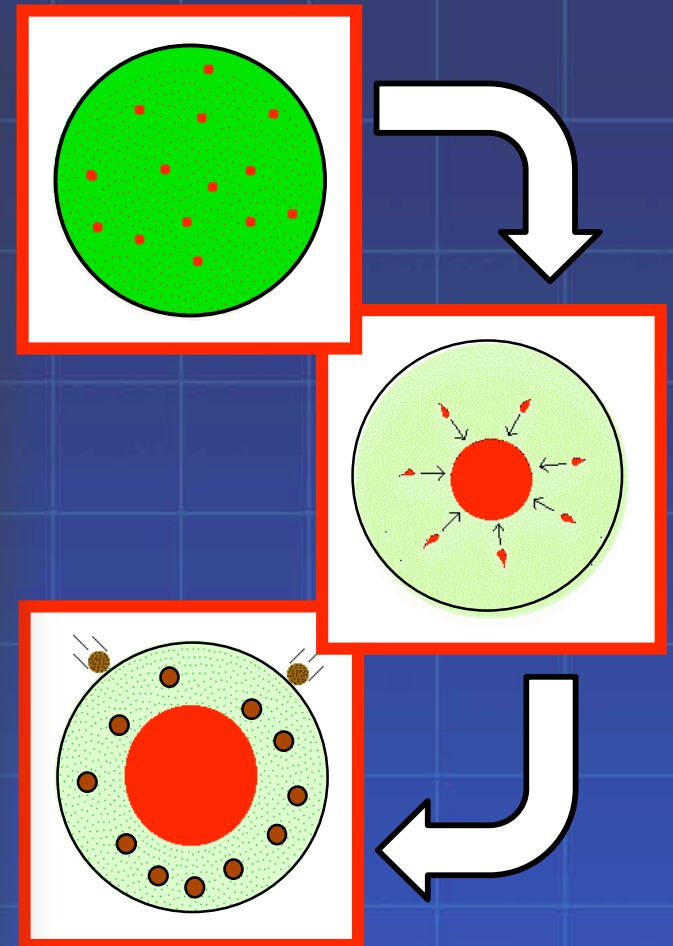
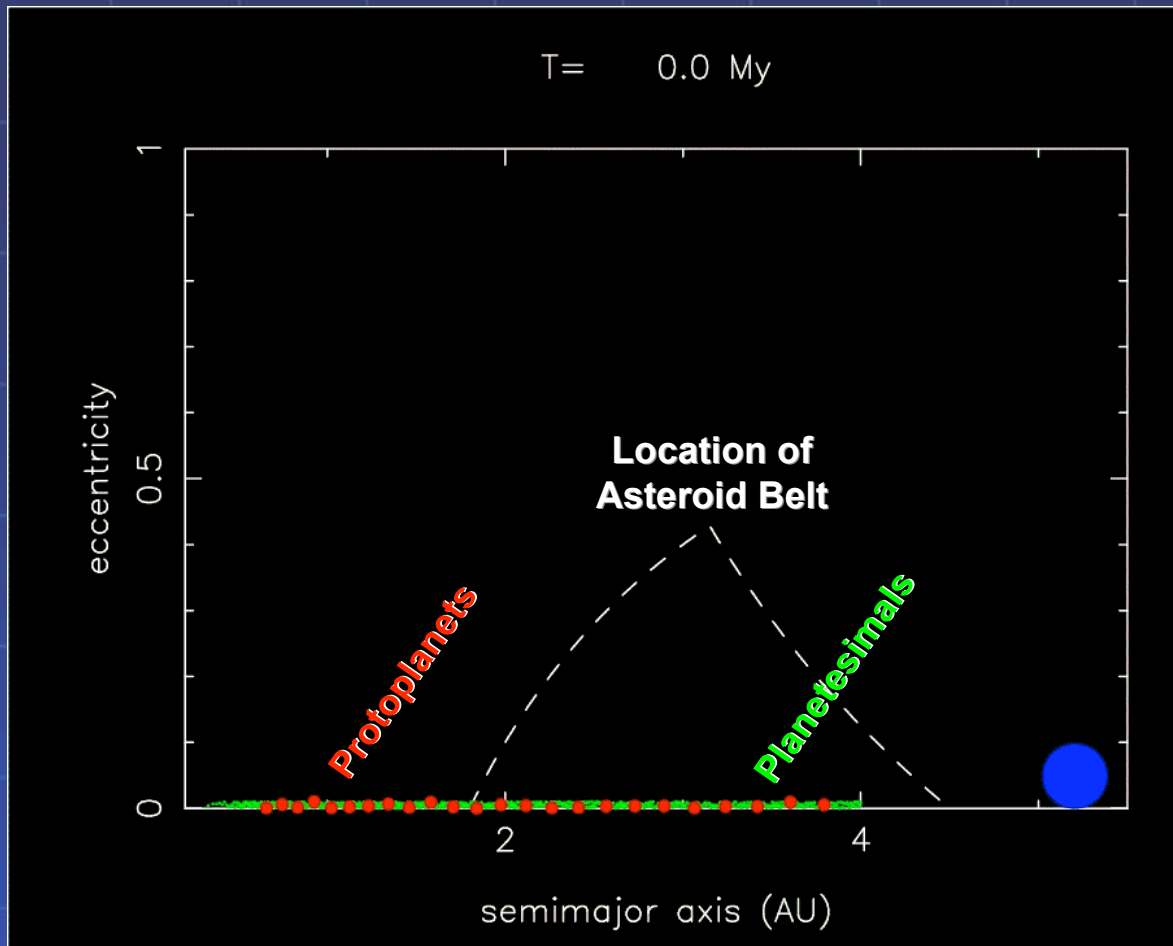
- The Moon is a factor of ~ 20 lower than the Earth in HSEs, but also has chondritic relative proportions. **Why?**

Part 3:

Late Accretion on the Earth and Moon

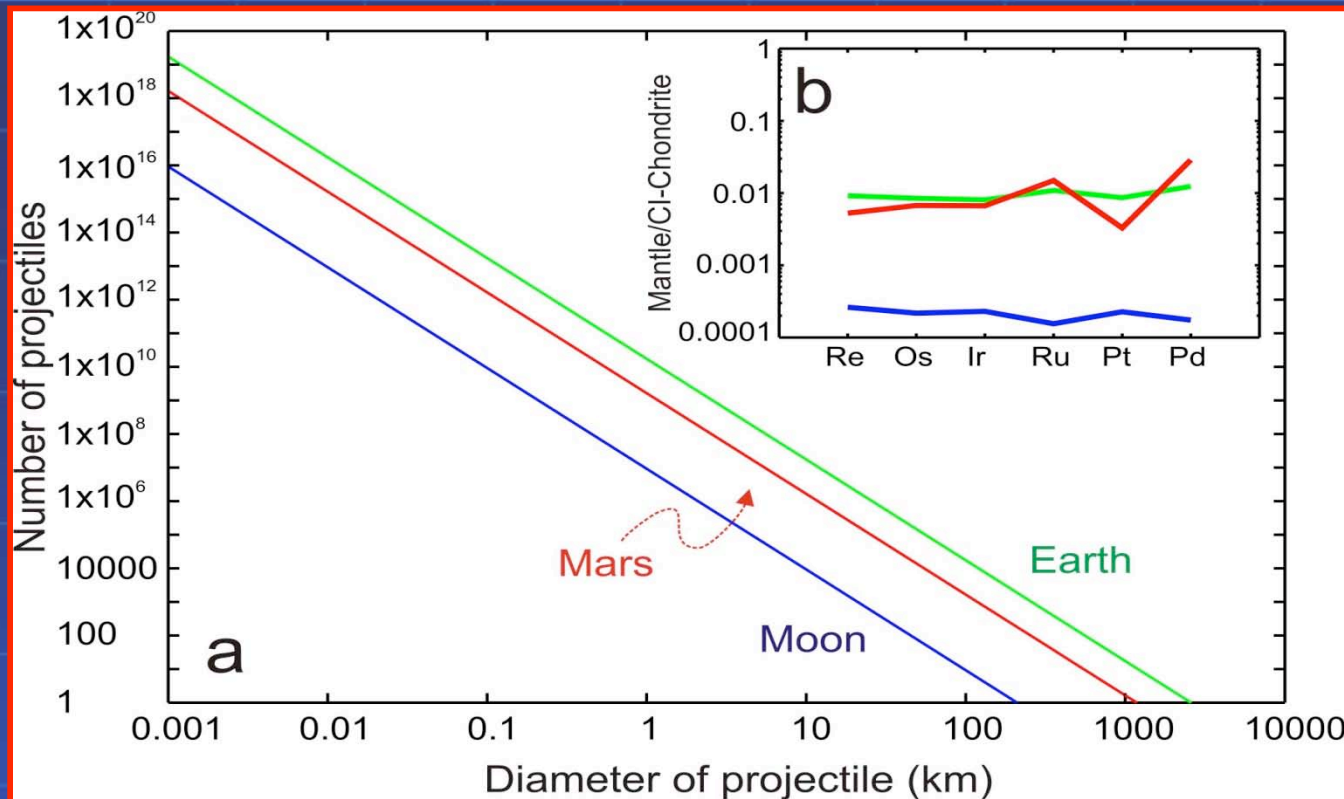


What is “Late Accretion”?



- Addition of “chondritic” material to the Earth during end stages of, or following core formation (Chou, 1978).

How Much Mass is Needed for Earth, Moon, Mars?

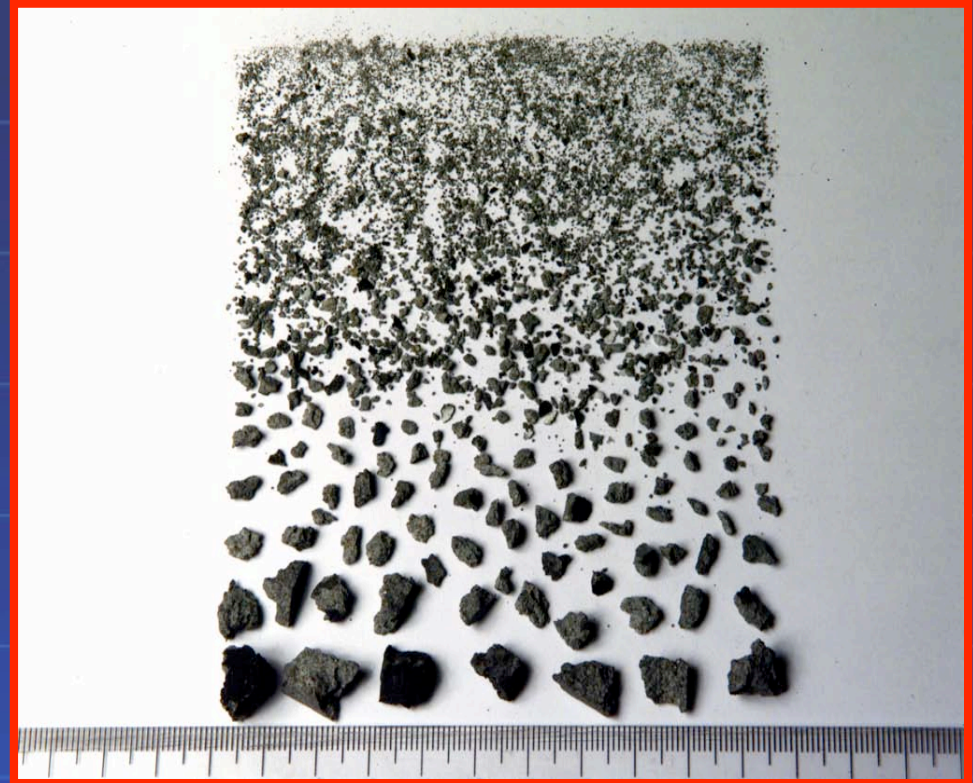
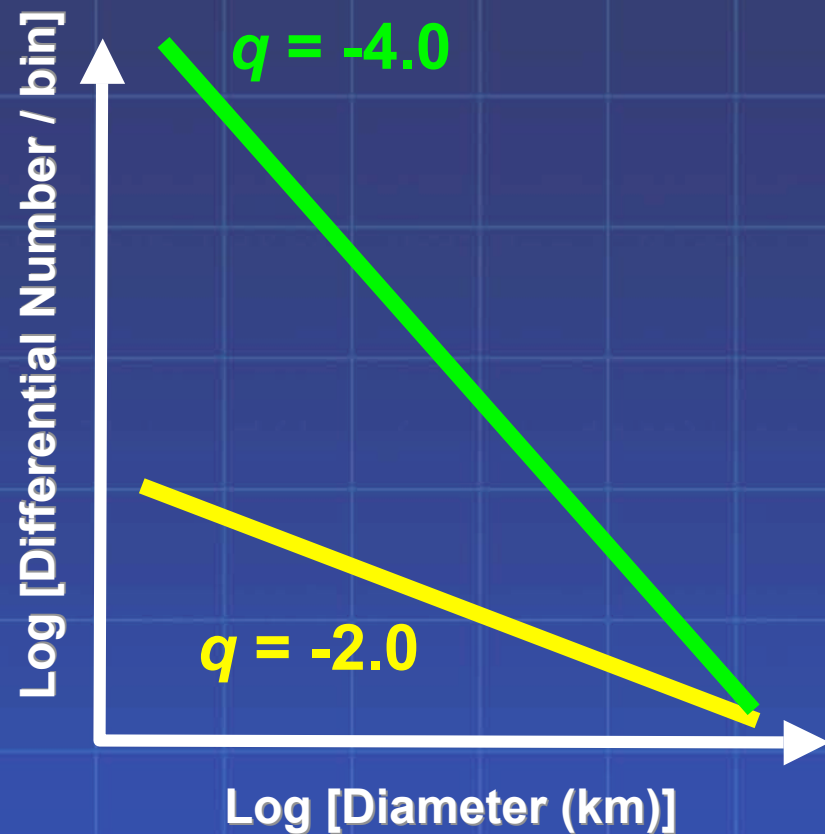


- Chondritic additions of $> 0.4\%$ of the Earth's mass are required to provide necessary HSEs.
- **We need a factor of 1,200 more mass for Earth than Moon!**

The Nature of Late Accretion

- The Earth/Moon see the same impacting population, with the impactors hitting in $\sim 20:1$ ratio.
- The input mass in Earth/Moon mantles need ratio of $\sim 1,200$.
- The Moon loses $\sim 40\%$ of projectile material upon impact. This moves Earth/Moon input mass ratio from $\sim 1,200$ to ~ 700 (e.g., Artemieva & Shuvalov 2008)

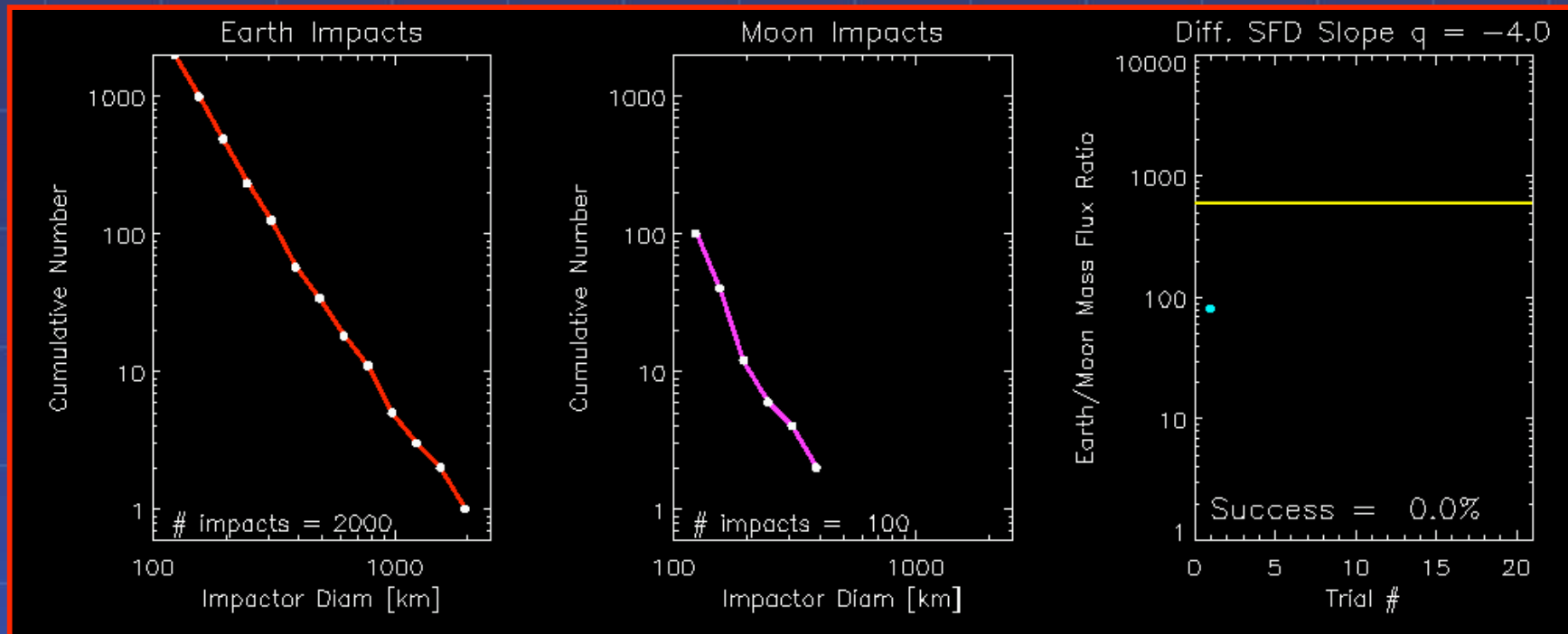
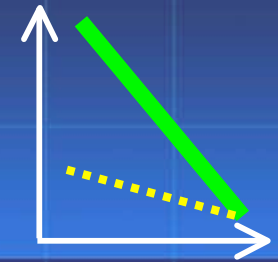
Testing Various Impacting Populations



- We decided to use a Monte Carlo code to test how different impacting populations affect the Earth and Moon.

Model #1

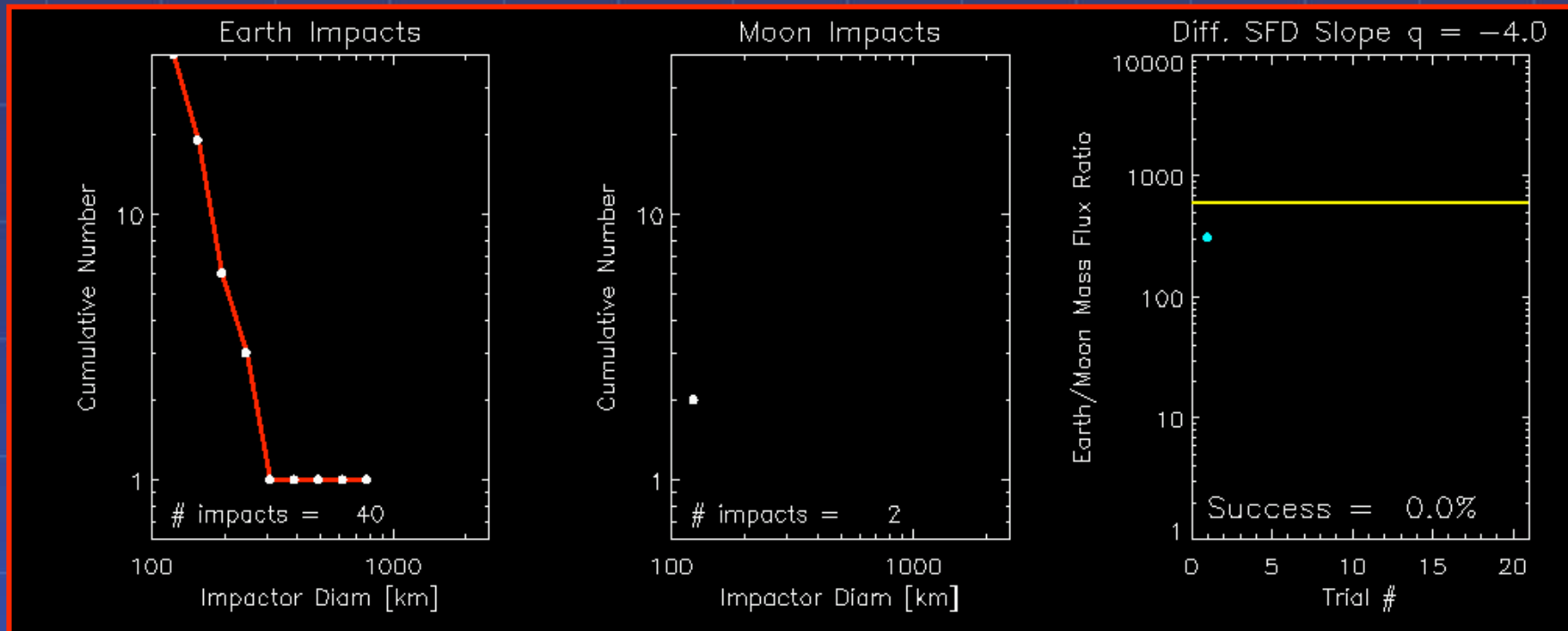
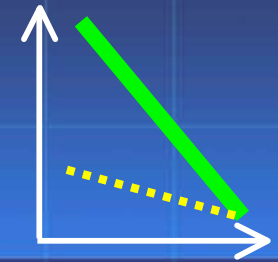
Many Impactors, Steep SFD



- Lots of tiny impactors ($q = -4$) does not yield a high input mass flux ratio.

Model #2

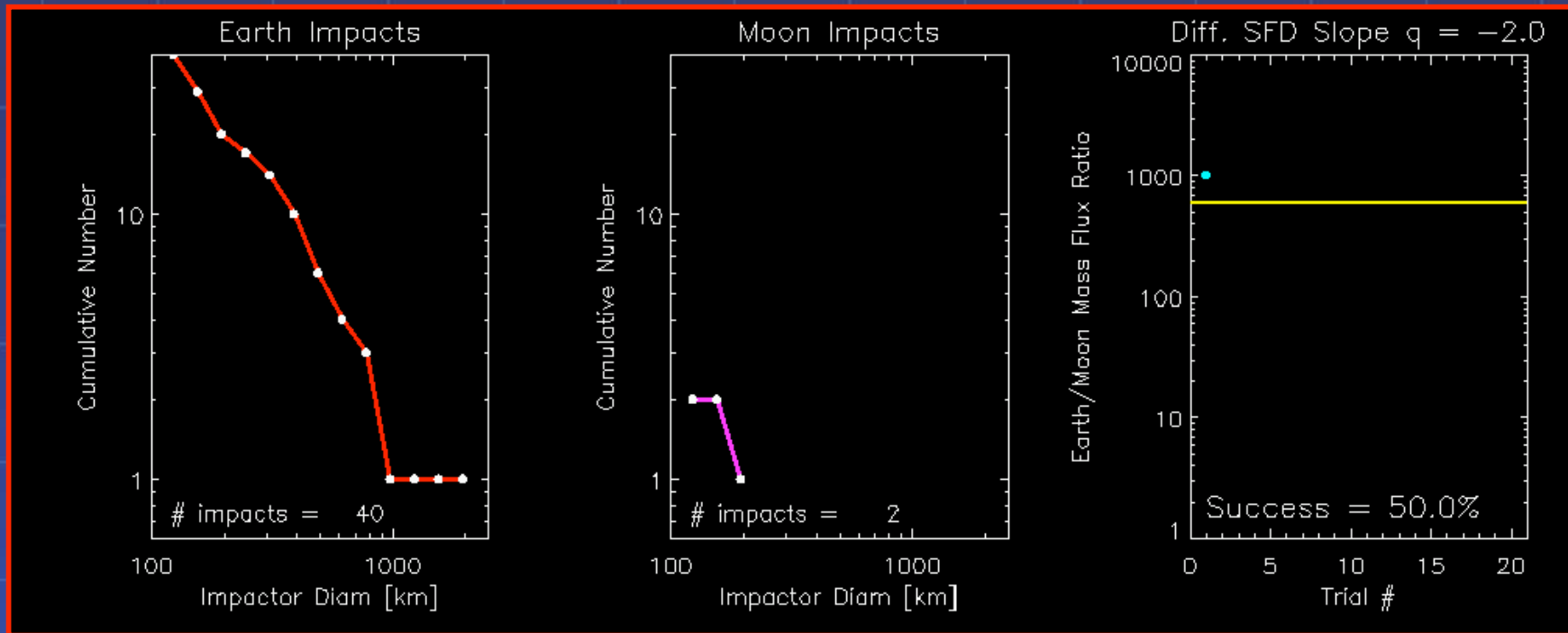
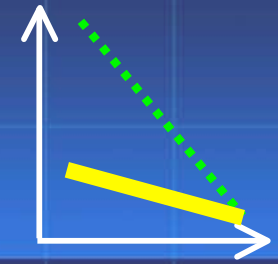
Few Impactors, Steep SFD



- Fewer impactors with steep size distribution ($q = -4$) also does not work. But...
- Stochastic variations yield mass ratios approaching ~700.

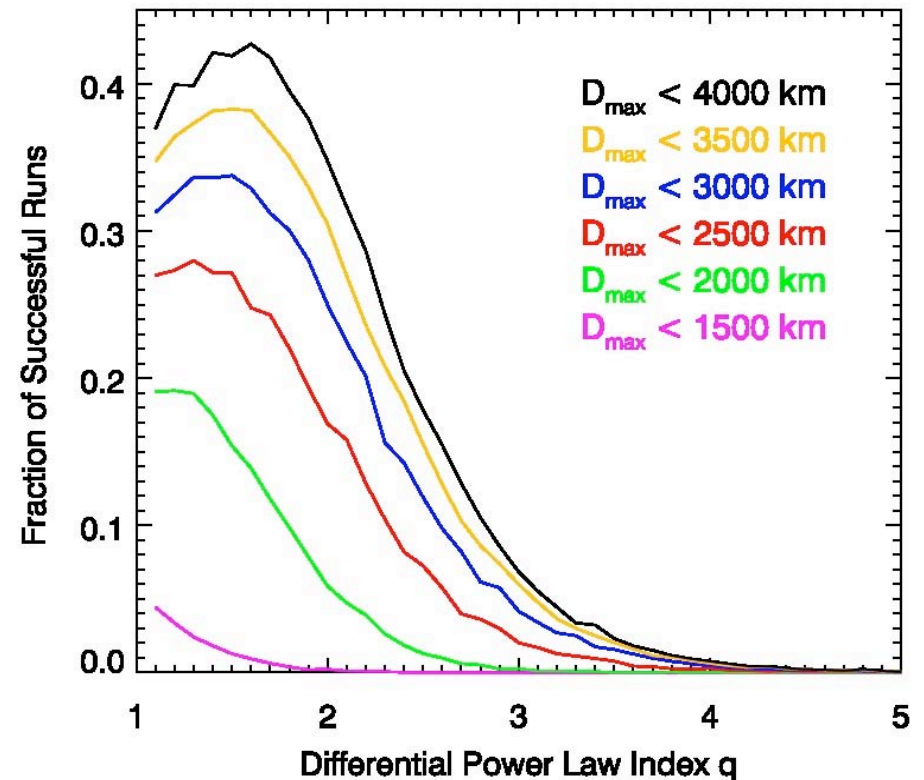
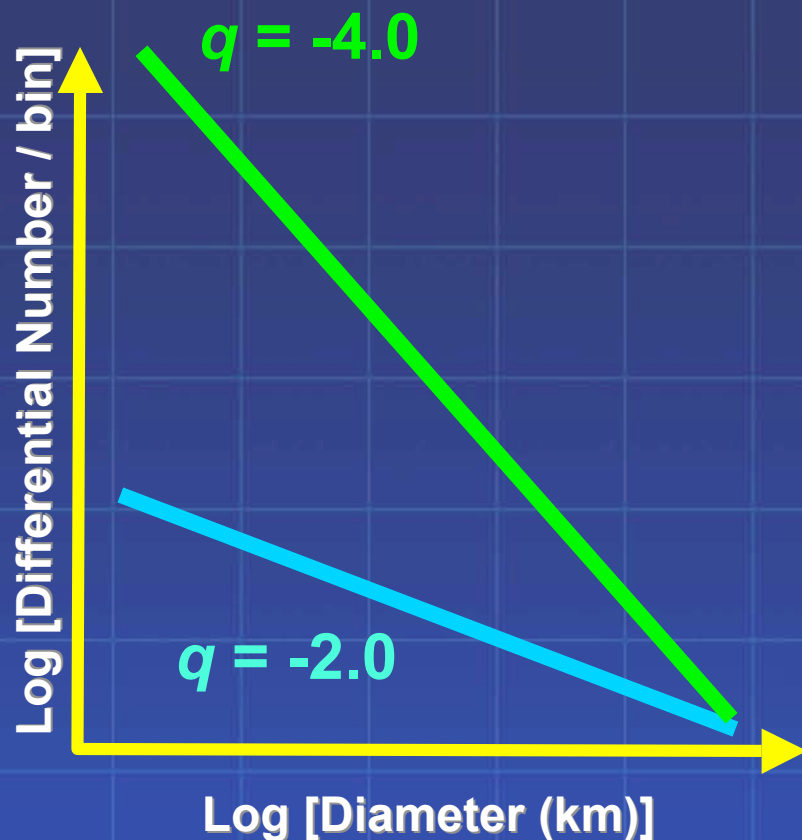
Monte Carlo Model #3

Few Impactors, Shallow SFD



- Few impactors with shallow size distribution ($q = -2$).
- On average, Earth hit by large impactors that miss Moon.
- Success rate approaches 25-30%

Late Accretion May Require Shallow Size Distributions

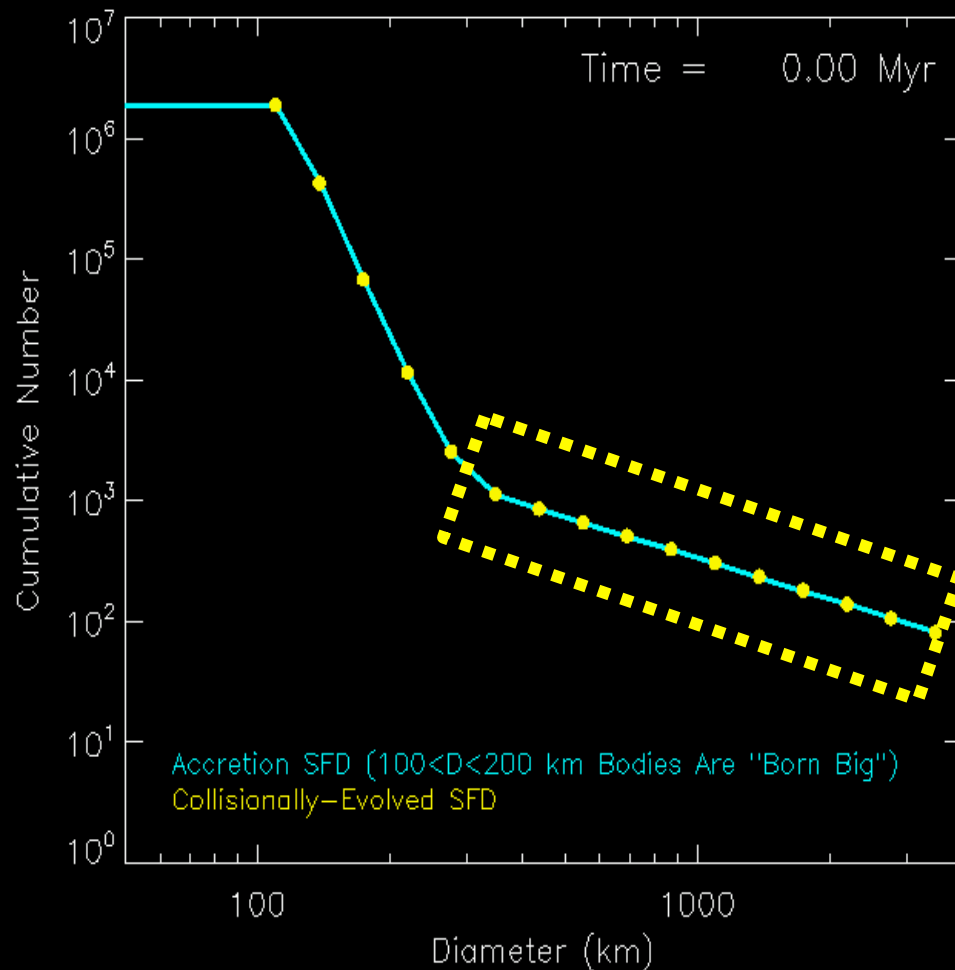


- We find that late accretion size distribution with most of their mass in largest bodies ($q < -2$) produces best results.

Part 4: **Evidence For a “Foot”** **in Late Accretion Populations**



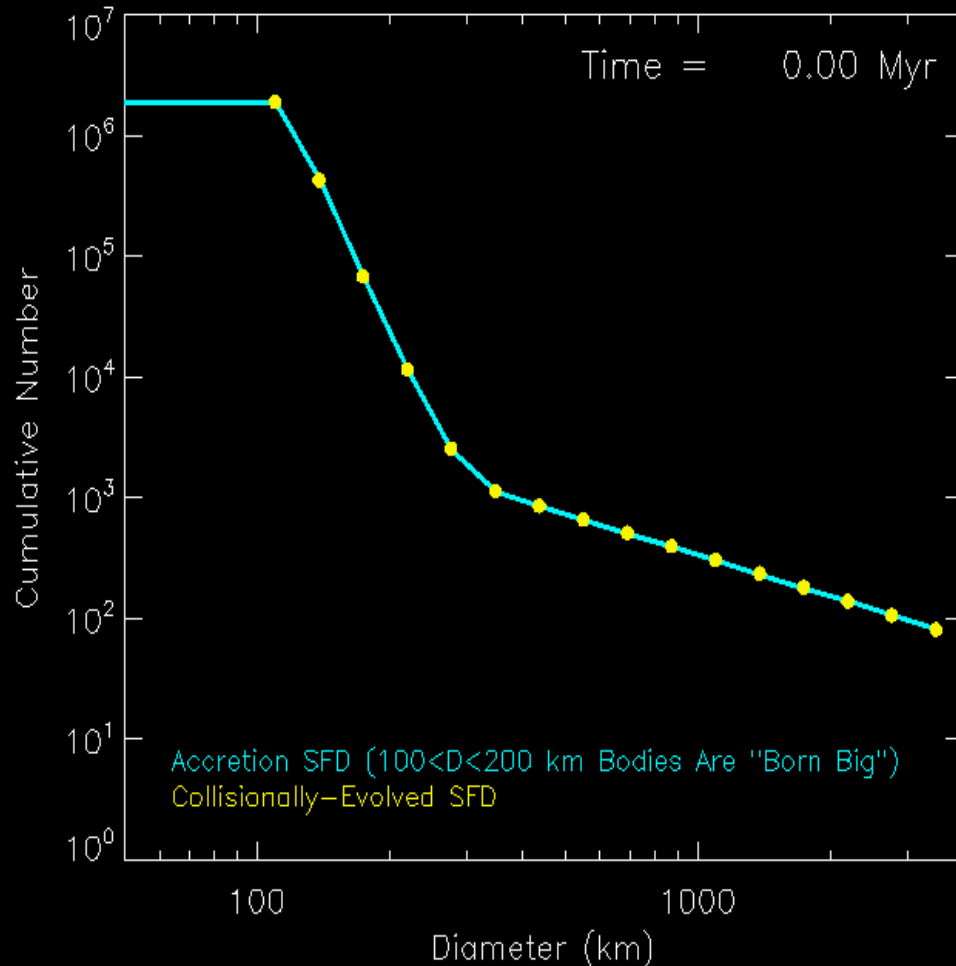
Evidence For A Shallow Late Accretion Population



- New planetesimal formation models make $D \sim 100$ km bodies.
- When inserted into accretion code, it produces a shallow "foot" for $D > 200$ km.
- The "foot" is $q \sim -2$.

Morbidelli, Bottke et al. (2009)

Evidence For A Shallow Late Accretion Population



■ Accretion SFD:

- The “foot” is resistant to collisional evolution for runs near 1 AU.

■ Inner main belt:

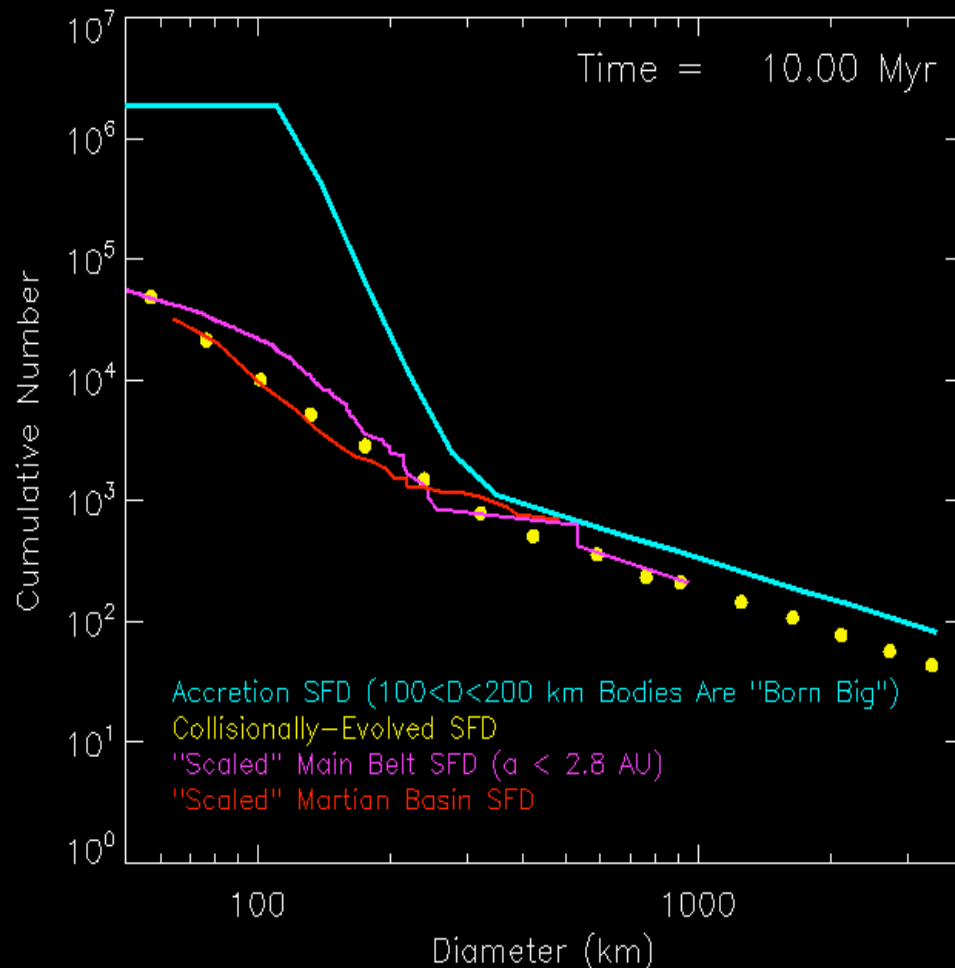
- A “foot” exists for $D > 250$ km asteroids.

■ Martian impact basins:

- A “foot” is seen when basins are changed to projectile diameters.

Morbidelli, Bottke et al. (2009); Frey et al. (2007)

Evidence For A Shallow Late Accretion Population



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Implications: Big Late Accretion Projectiles

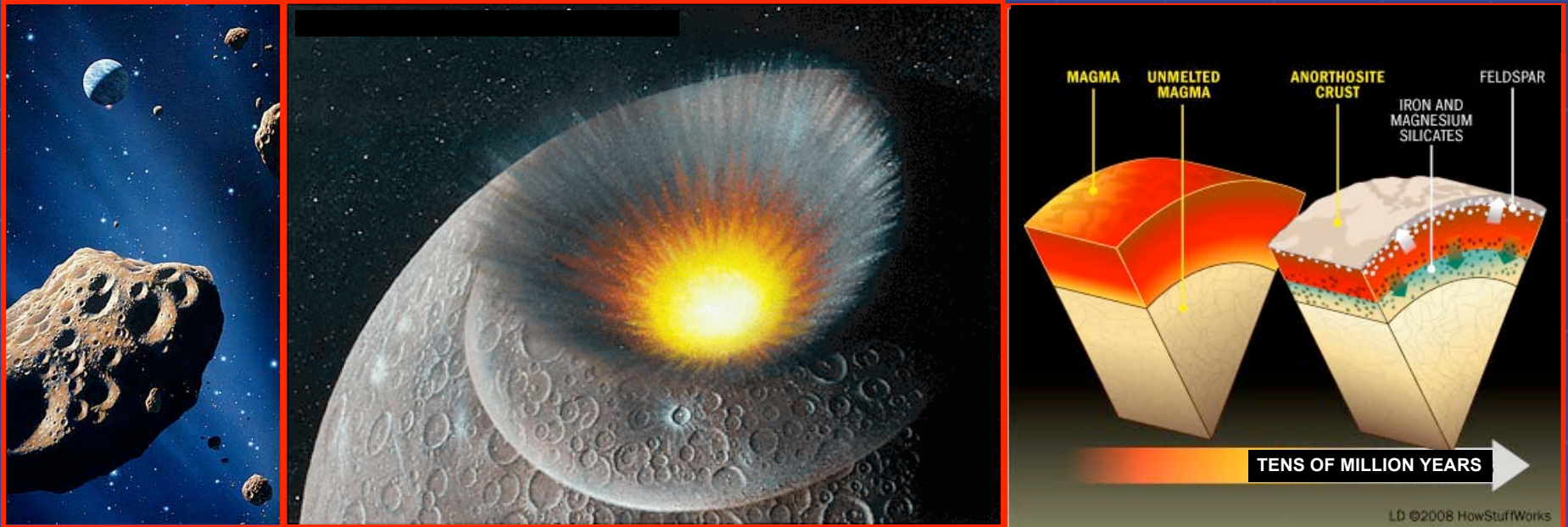
■ Diameter of largest late accretion projectiles to strike Earth, Moon, and Mars:

Earth	Moon	Mars
2500-3100 km	250-280 km	1500-1800 km

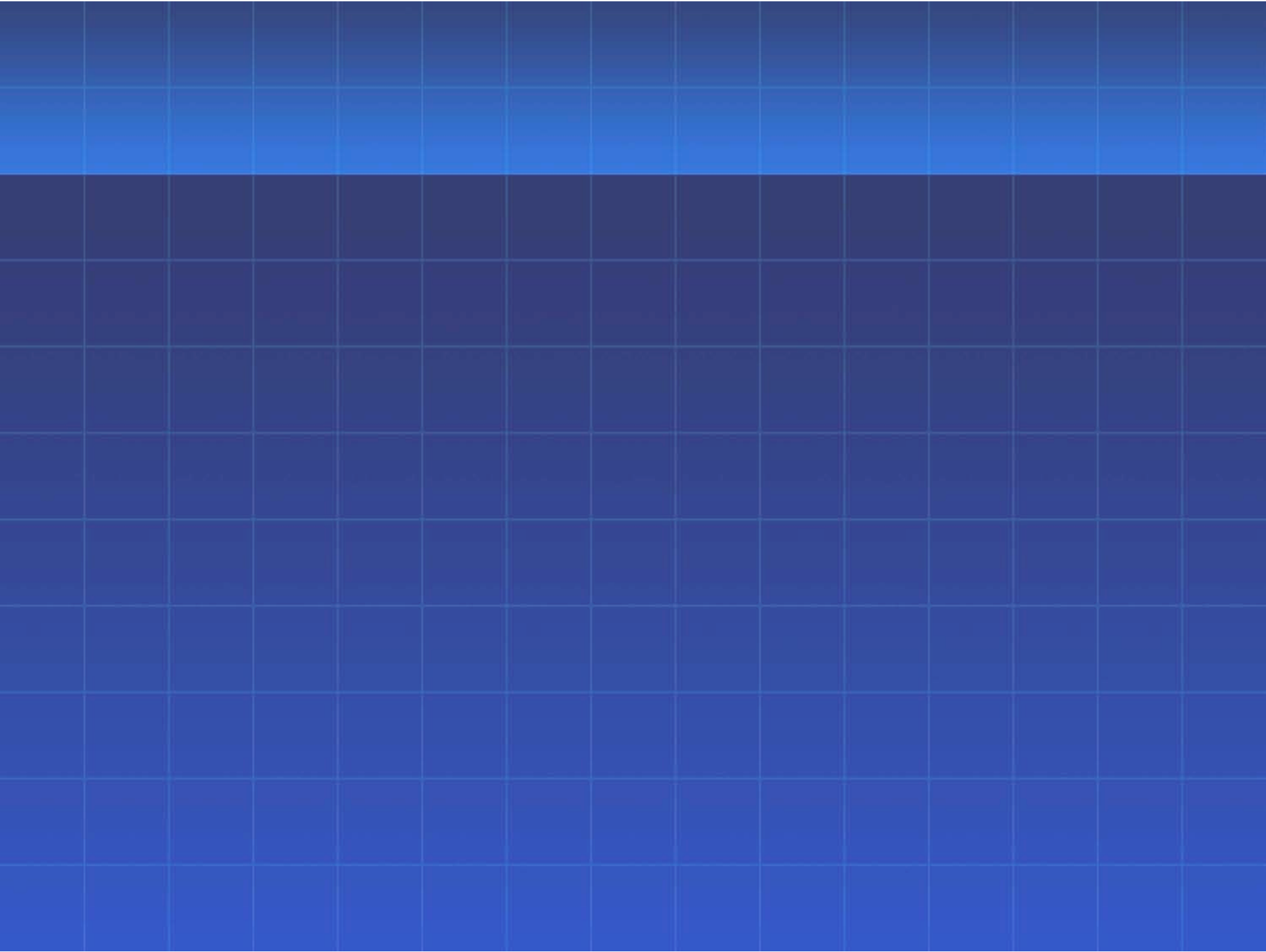
- Impact modifies Earth's obliquity by $\sim 10^\circ$. Can this explain the inclination of Moon's orbit?
- Lunar impactor large enough to produce South-Pole Aitken basin (or possibly Procellarum basin).
- Martian impactor is the right size to make gigantic Borealis basin.

Implications:

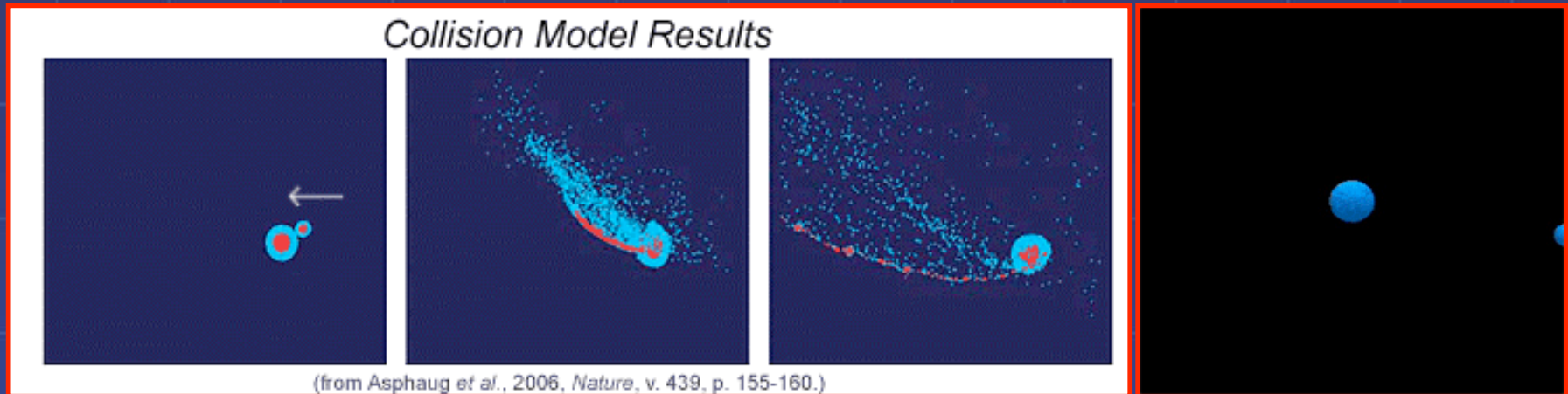
Did Lunar Mantle Water Come from Late Accretion?



- Assume the Moon was hit during magma ocean phase:
 - $D = 250\text{-}280$ km projectile
 - Assume it had 0.1% water and was mixed into lunar mantle between depths of 100-500 km.
- This yields a 1-3 ppm wt% water, the same values estimated from lunar apatites (McCubbin et al. 2010).

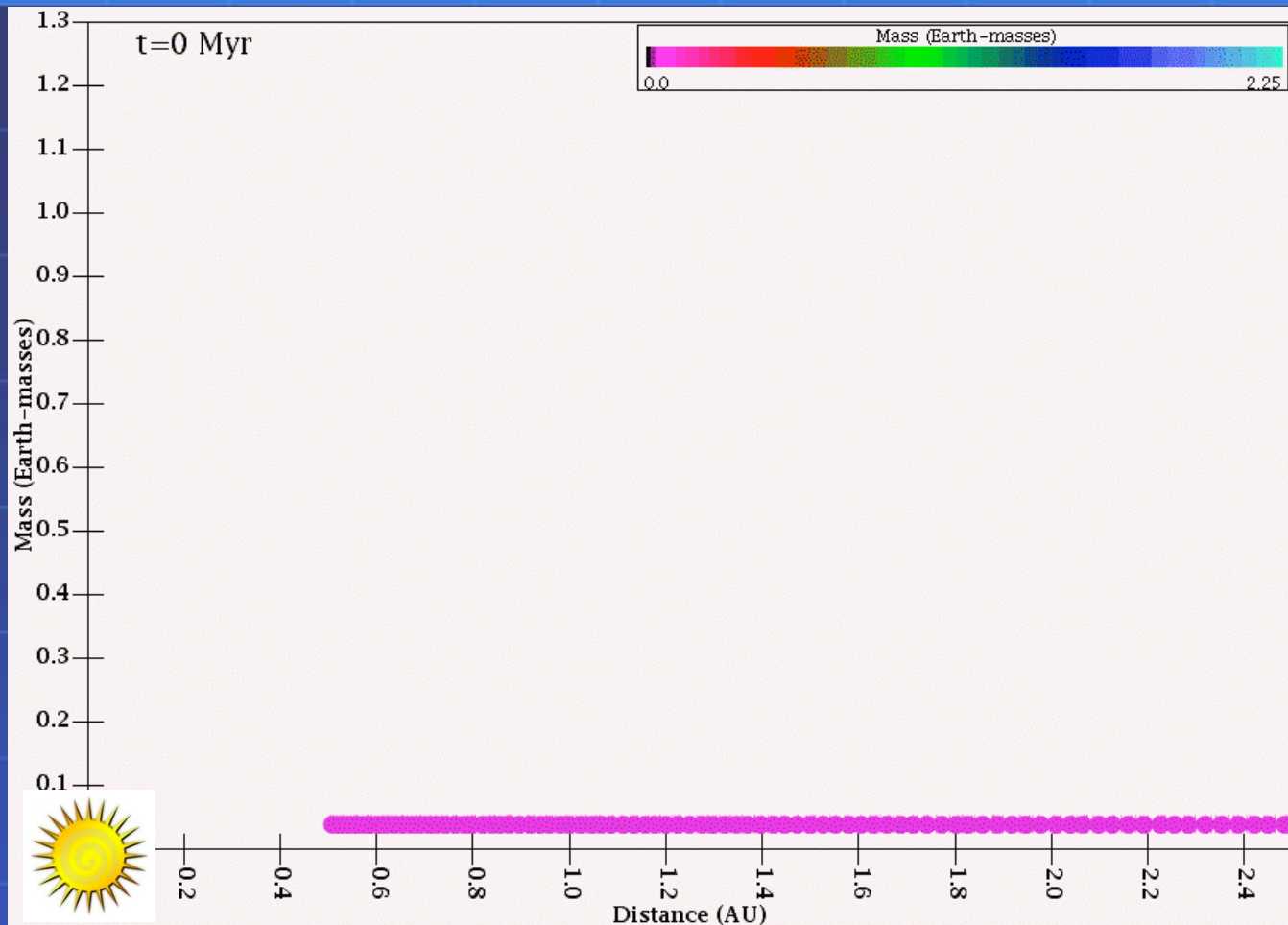


Late Accretion on the Earth: A Case of “Hit and Nearly Run”



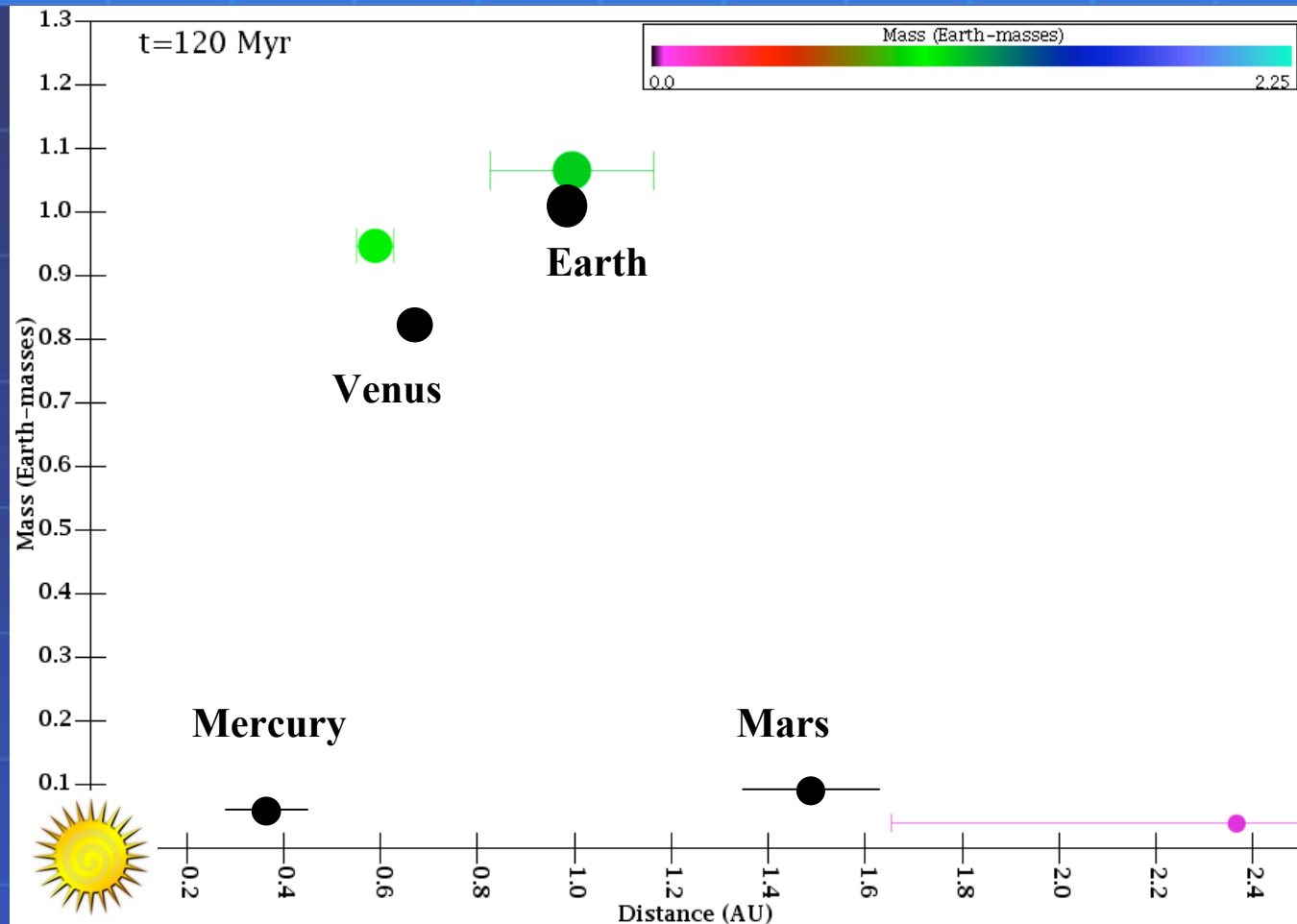
- $D = 2500\text{-}3200$ km impactors on Earth should act like “hit and nearly run” collisions.
- Most of the projectile’s core escapes immediate accretion but the core fragments are eventually re-agglomerated.
- The iron and HSEs possibly emulsify into mantle immediately or are slowly incorporated into mantle via plate tectonics.

Simulated Planet Growth



- Starting with several hundred “mini-planets”, collisions cause bodies to merge and form big planets!

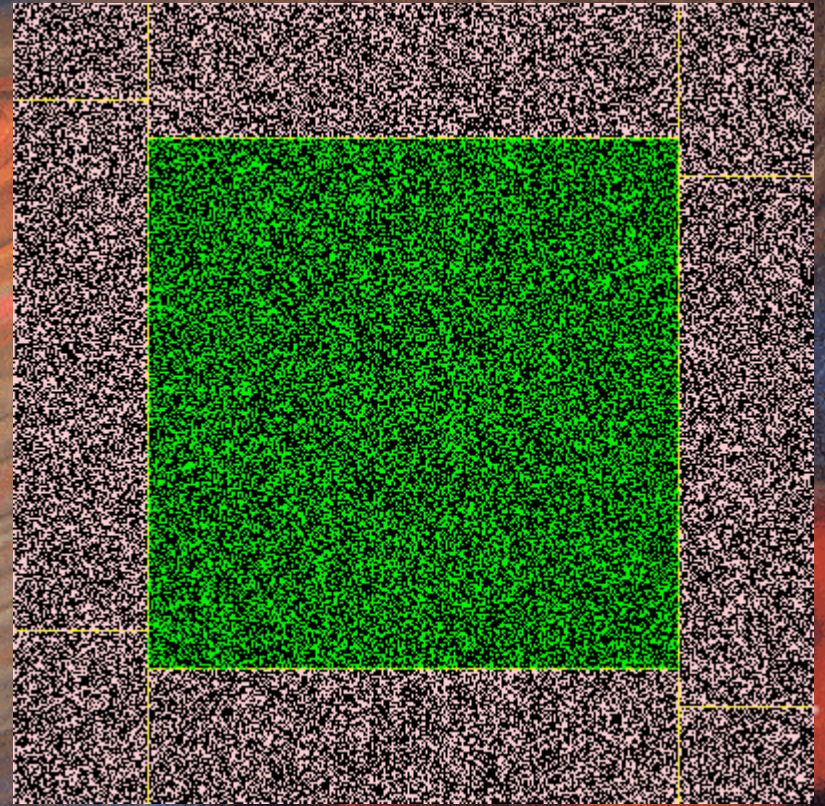
Simulated Planet Growth



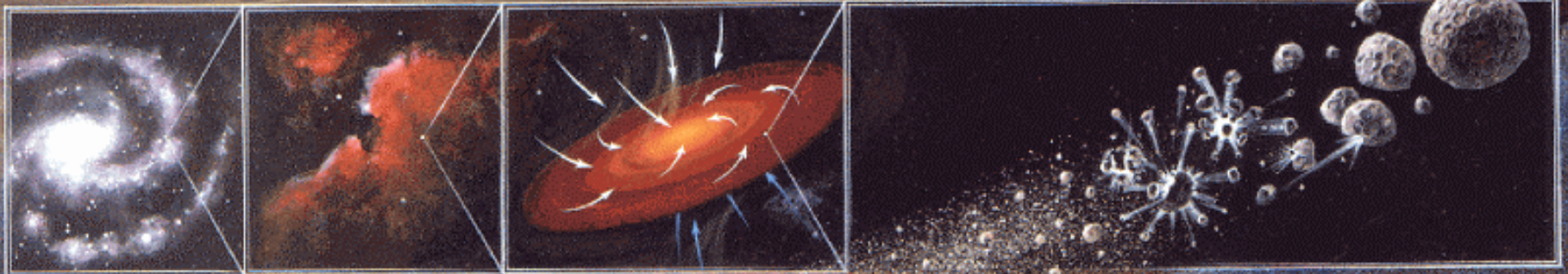
■ In the end, we end up with model planets like our own.

Planetesimal Formation

- Newly-formed Sun surrounded by an orbiting disk of gas and dust.
- Disk particles come together by gravity. Collisions make larger and larger objects by “accretion”.



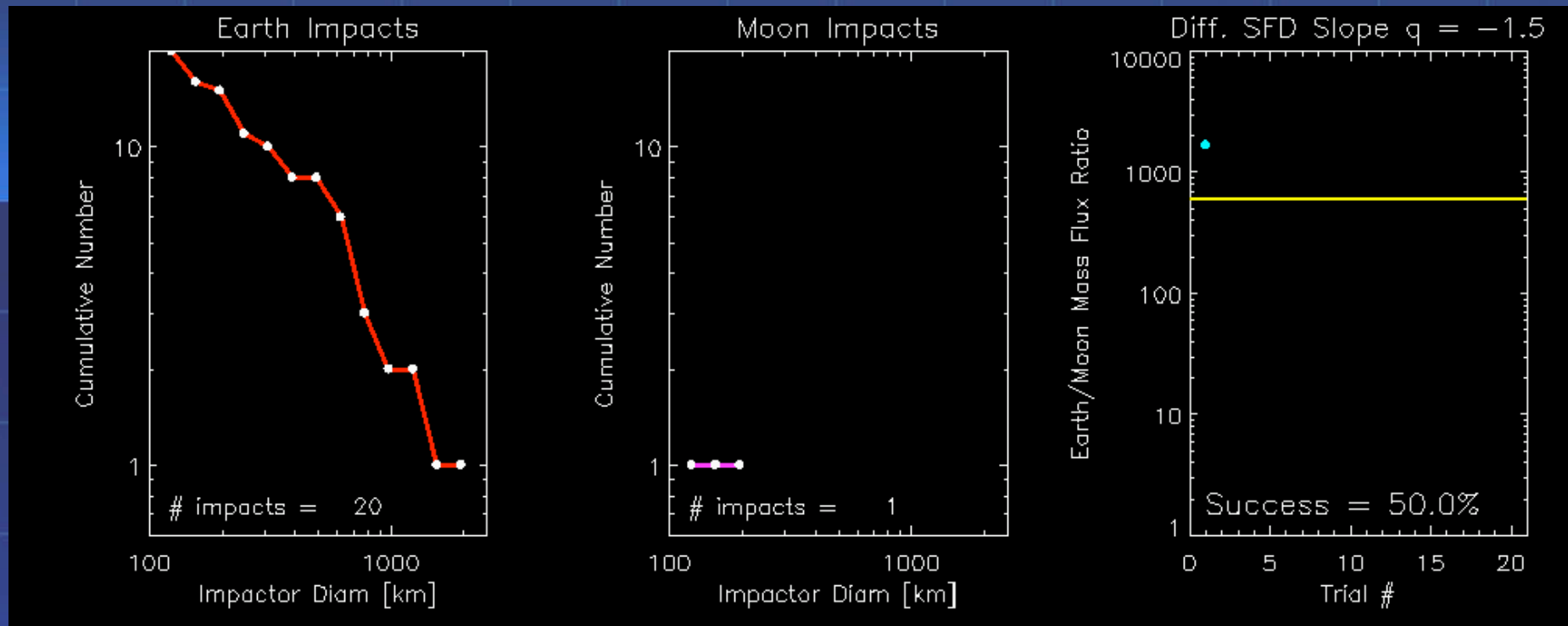
Animation from Tanga et al. (2003)



More Late Accretion Constraints



- “Pristine” lunar rocks have very low HSEs and probably dominate lunar crust. This suggests crust is unlikely to be a major reservoir of HSE.
- The oldest known sample of the lunar crust formed ~100 My after CAI formation (4.46 Ga).
- Late accretion impactors need to hit within a few tens of My after Moon formation to supply HSEs.

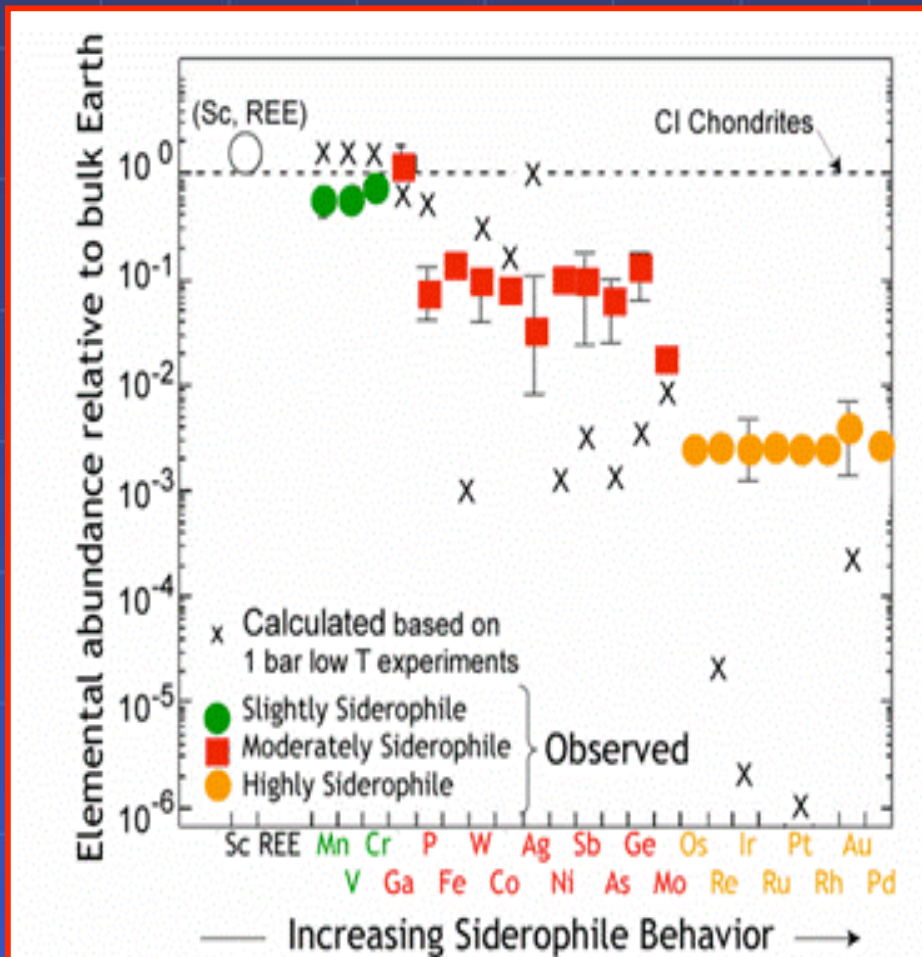


- **Model 4:**

- **Few impactors, with a shallow size distribution ($q = 1.5$).**

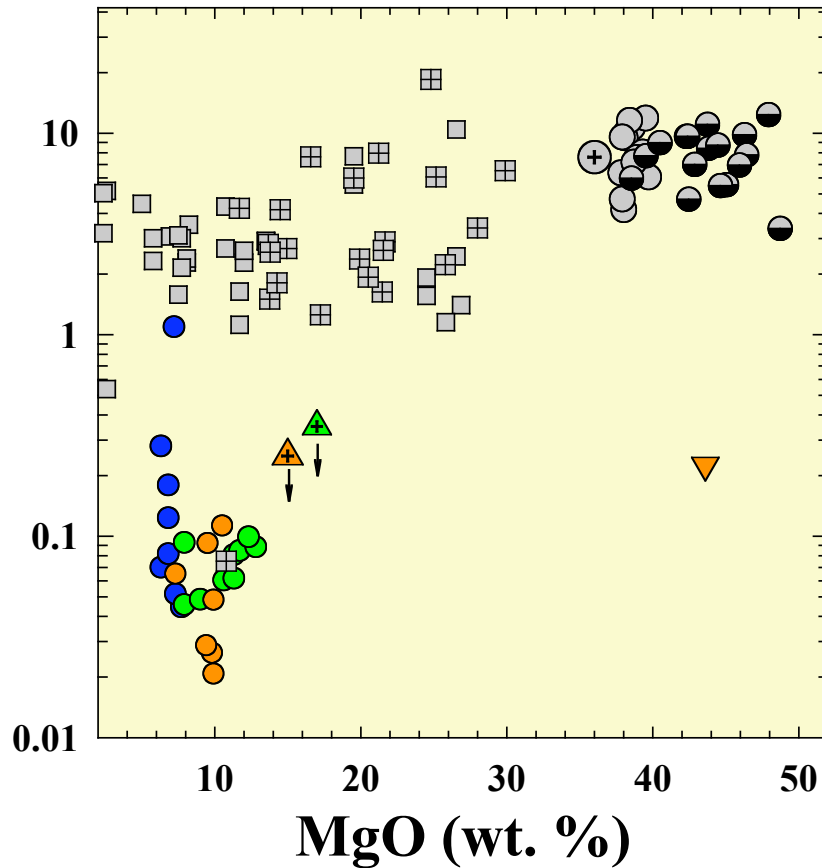
- **Success rate of $>50\%$**

Highly Siderophile Elements in Earth's Mantle



- Mantle peridotites indicate Earth's ancient mantle was only depleted in HSEs by factor of ~200 compared to chondrites.
- If HSEs are mixed throughout mantle, **chondritic additions of ~0.4% of the Earth's mass** are required to provide necessary HSEs.

Highly Siderophile Elements in Lunar Mantle



- HSE abundances are apparently very low.
- HSE versus MgO plots consistent with **>20 times depletion relative to terrestrial primitive upper mantle.**

Walker et al. (2004); Day et al. (2007)

Take Away Message



- Big events on Earth and Moon are linked in time.
- The Earth and Moon have similar HSE signatures.
- The mass added to Earth was higher by factor of ~1,200!
- How do we get this?